

Development of cloud algorithm for GEMS: Synthetic cloud data

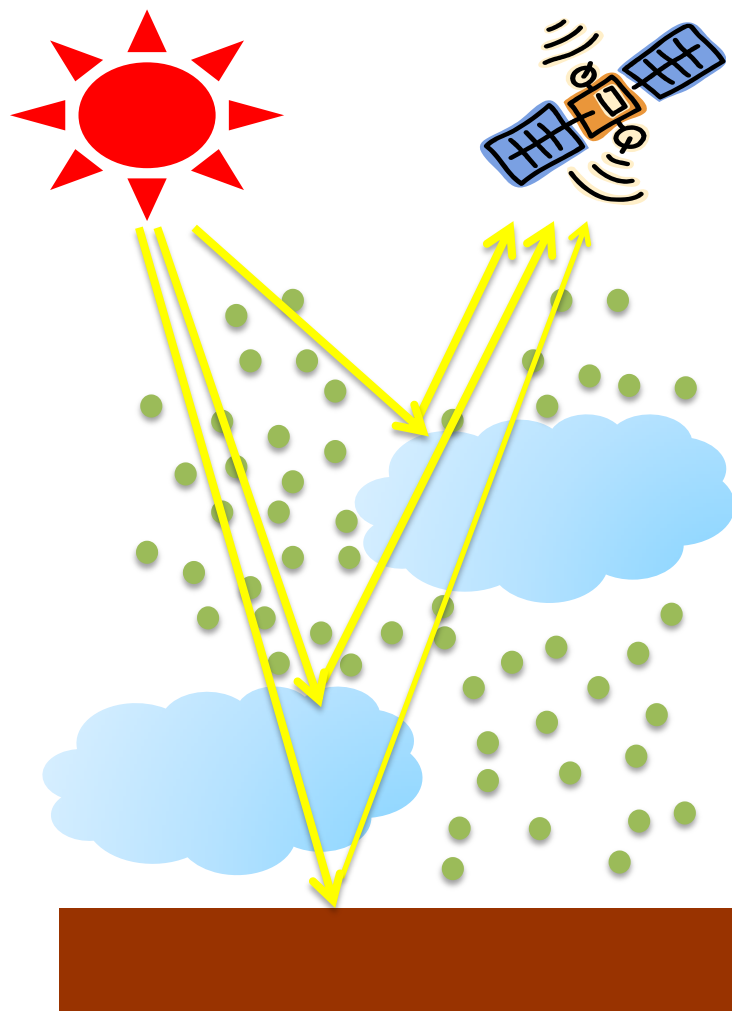
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Introduction

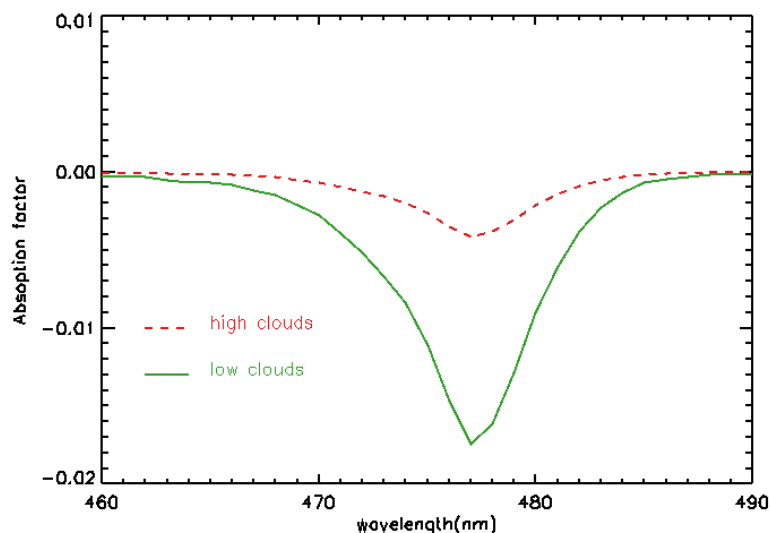


- Cloud contamination causes
 - ➔ errors in air mass factor
 - ➔ errors in gas/aerosol retrievals
- Accurate cloud data will serve to
 - ➔ better **cloud screening** for gas/aerosol retrieval
 - ➔ better **cloudy-sky gas/aerosol retrievals***

**Vasilkov et al. (2004) Improving total column ozone retrievals by using cloud pressures derived from Raman scattering in the UV"*

Theoretical background for cloud pressure retrieval

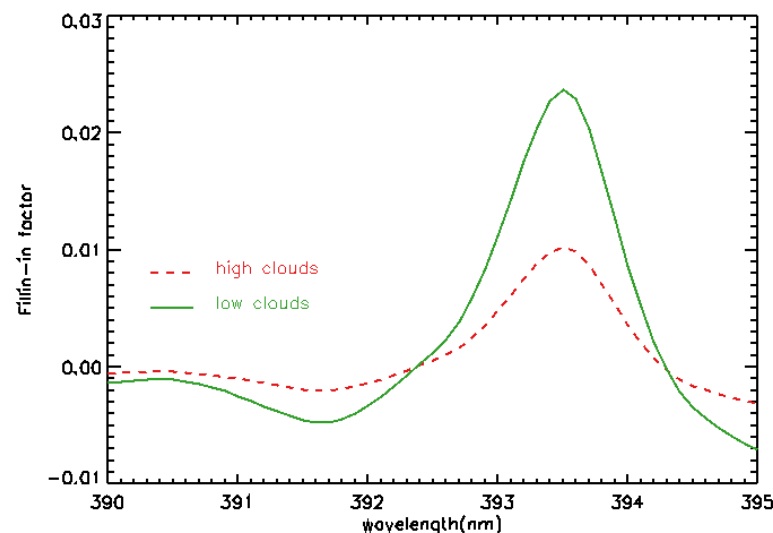
O₂-O₂ absorption (477 nm)



$$A(\lambda) = [R_m(\lambda) - R_f(\lambda)]/R_f(\lambda)$$

- $R_m(\lambda)$: the measured reflectance
- $R_f(\lambda)$: the calculated reflectance using without O₂-O₂ absorption

Rotational Raman scattering (393.6 nm)



$$k(\lambda) = [I_m(\lambda) - I_R(\lambda)]/I_R(\lambda)$$

- $I_m(\lambda)$: the measured radiance
- $I_R(\lambda)$: the calculated radiance using only Rayleigh scattering

The cloud retrieval algorithm and its validation problems

Input Data

- Radiance and spectral reflectance ($R = \pi L/E$) of 300~500 nm
- Location (longitude, latitude, local time)
- Calculated scattering angle from location data

Using O₂-O₂ absorption band at 477 nm

Calculating cloud effective fraction

$$c = \frac{(R_{meas} - R_{clear})}{(R_{cloud} - R_{clear})}$$

- R_{meas} : the measured reflectance
- R_{clear} : the reflectance for clear sky from surface albedo database
- R_{cloud} : the calculated reflectance of a completed cloudy pixel

Calculating cloud pressure

1. DOAS fit : to determine the continuum reflectance and the slant column density O₂-O₂
2. A look-up table : to determine the altitude, z_L
3. The cloud altitude z_L is translated to the cloud pressure, $p_L(O_2)$

Using Ring effect at 393.6 nm

Calculating cloud effective fraction

$$f = \frac{(I_{clear} - I_{meas})}{(I_{clear} - I_{cloud})}$$

- I_{meas} : the measured TOA radiance
- I_{clear} : the calculated radiance for clear sky
- I_{cloud} : the calculated radiance of a completed cloudy pixel

Calculating cloud pressure

1. Filling-in factor, k : to quantify the RRS (Rotational Raman Scattering) effect on radiance

$$k(\lambda) = [I_m(\lambda) - I_R(\lambda)]/I_R(\lambda)$$
 - $I_m(\lambda)$: the measured radiance
 - $I_R(\lambda)$: the calculated radiance using only Rayleigh scattering
2. A look-up table : to determine the altitude, z_L
3. The cloud altitude z_L is translated to the cloud pressure, $p_L(RRS)$

Output data

- Cloud effective fraction
- Cloud pressure

Is the cloud retrieval algorithm correctly working?

- To validate the cloud retrieval algorithm, we need cloud data, which is physically consistent with radiance data.

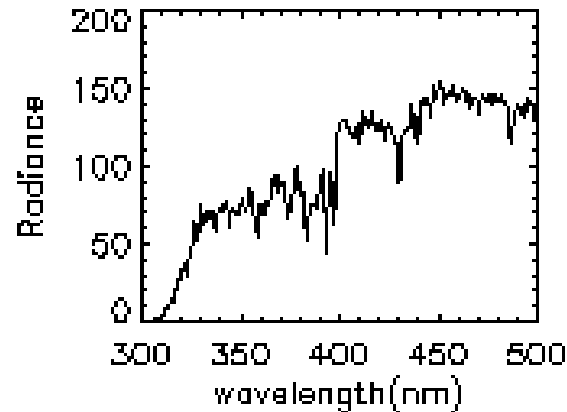
(Problem 1)

Every cloud data has observational errors.

(Problem 2)

No geostationary spectral UV/VIS radiance observations exist.

Concept of synthetic (or proxy) data



**Forward RT
simulation**

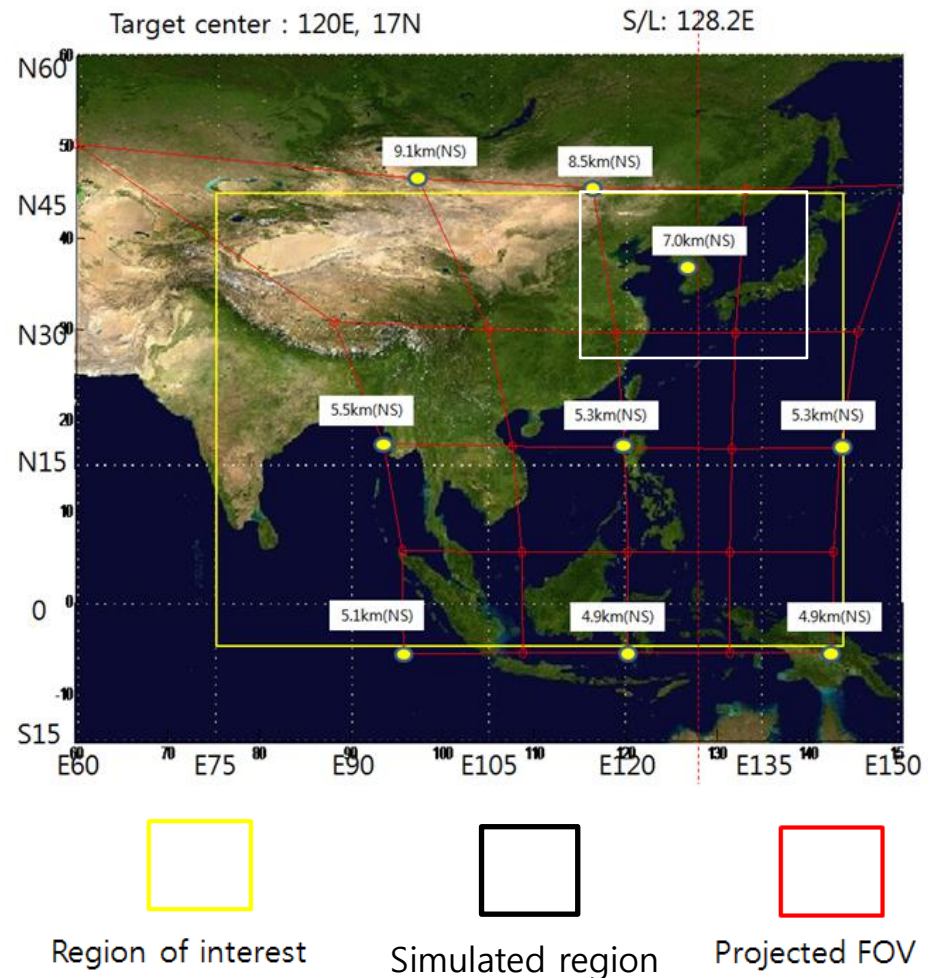
Retrievals by our
cloud algorithm



Known (WRF) cloud profiles

WRF (Weather Research & Forecasting Model) Data

Area: 25° – 45° N, 110° – 140° E

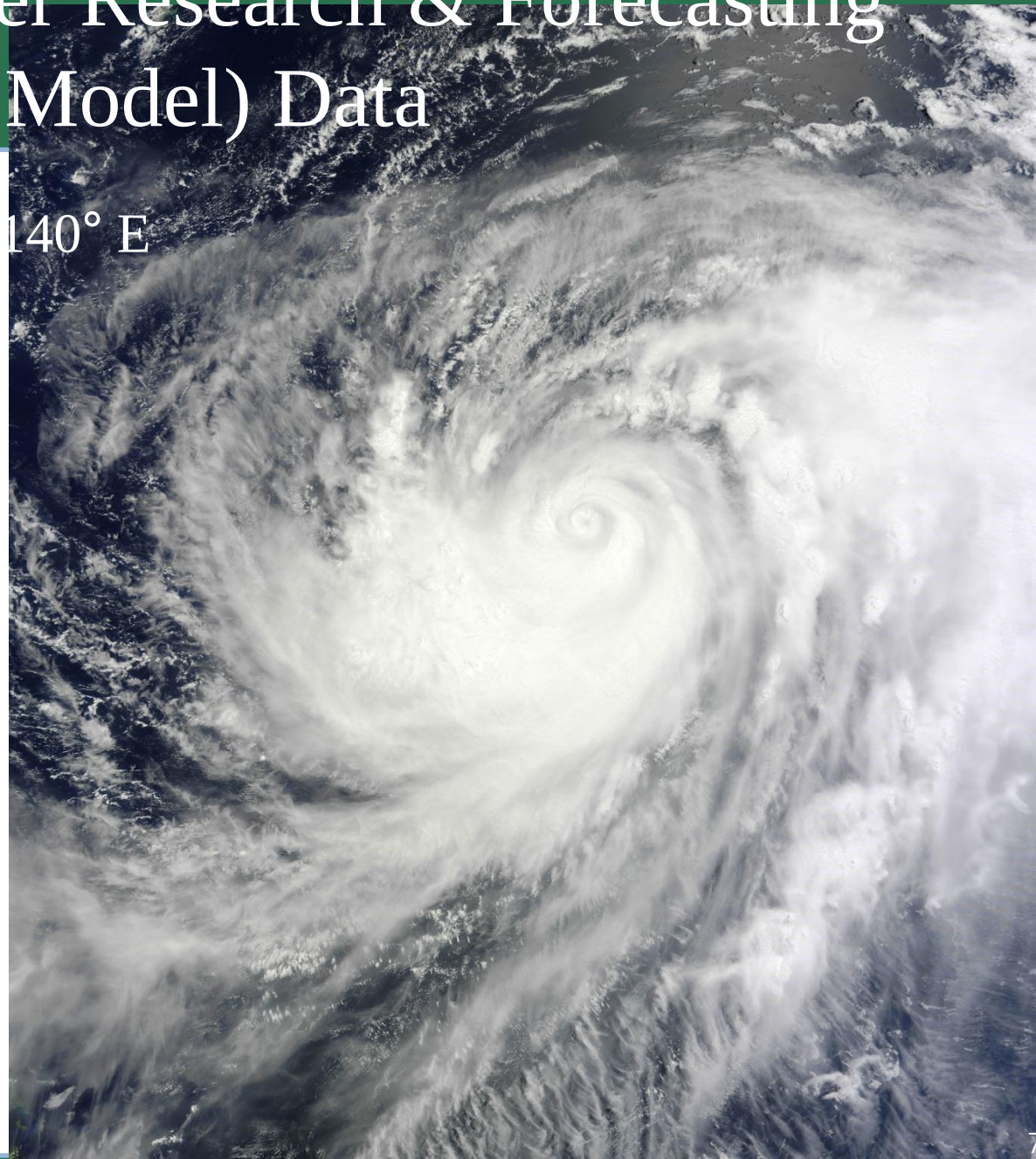


WRF (Weather Research & Forecasting Model) Data

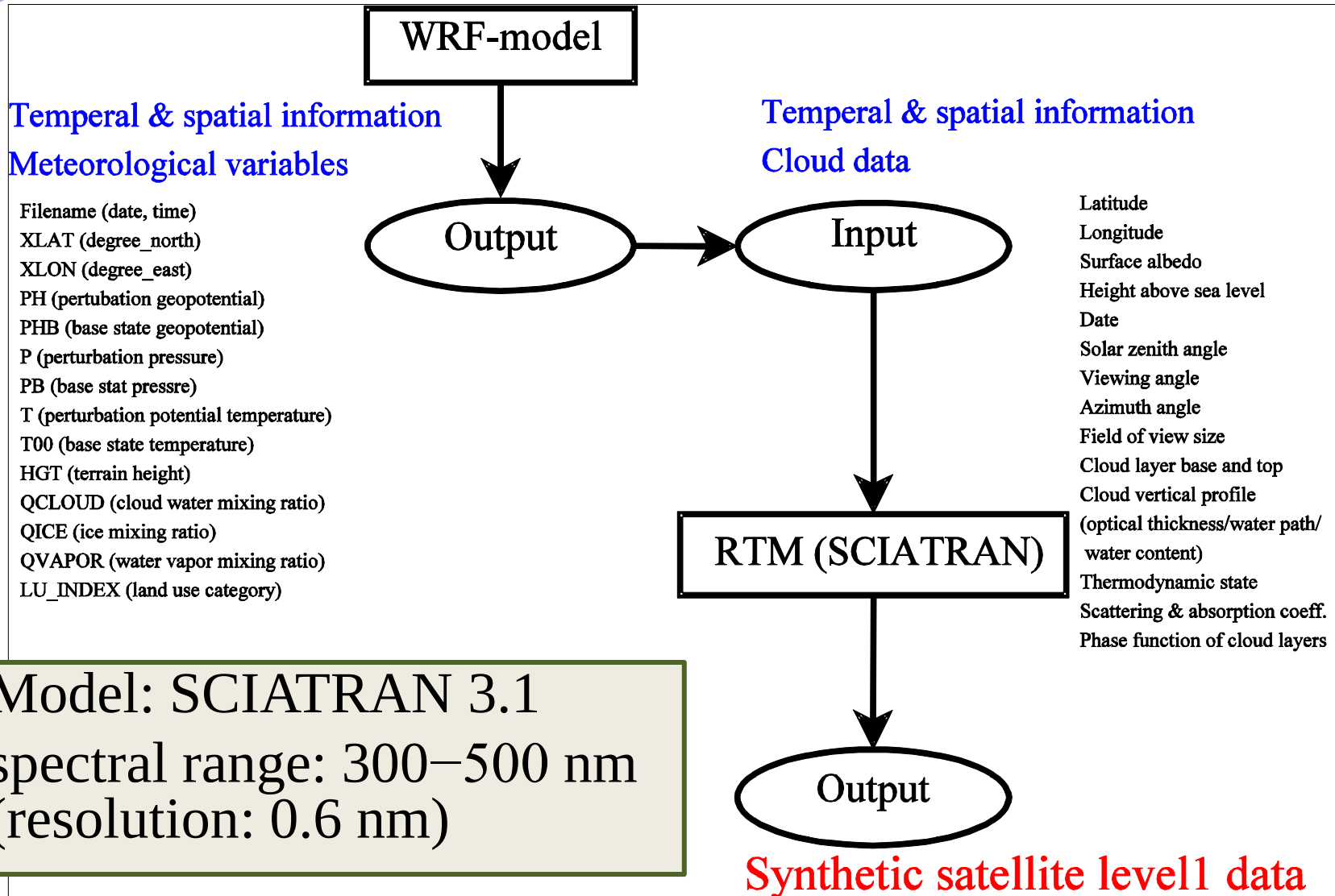
Area: 25° – 45° N, 110° – 140° E

Period: Aug. 6–7, 2011

(Typhoon Muifa)



Radiative transfer simulation



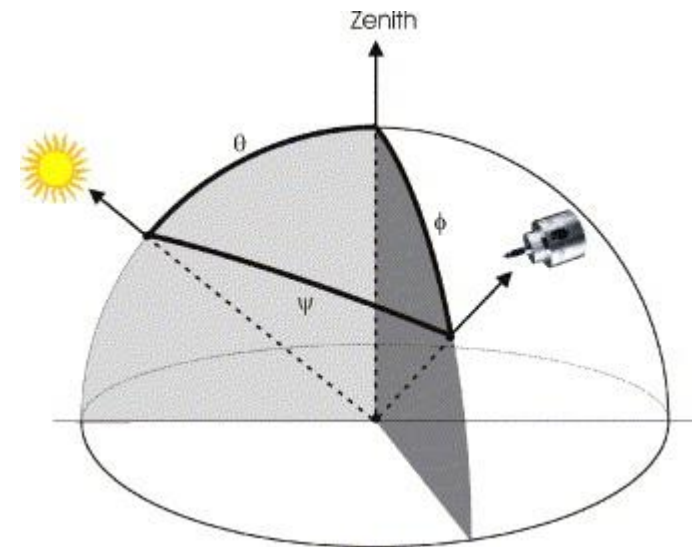
Assumptions for RT simulation

- Land use type
 - sand, snow, water, vegetation, or soil
- Cloud
 - All clouds are composed of 1-layer and **vertically homogeneous** (due to vertical inhomogeneity errors in SCIATRAN)
 - Size of liquid particle: $12\ \mu\text{m}$
 - Size of ice particle: $100\ \mu\text{m}$
 - Shape of ice particle:
“fractal” (SCIATRAN 3.1)



Angle geometry

- RT simulations are performed with due considerations of the angle geometry.
 - *Solar zenith angle* = $F(\text{lat}, \text{lon}, \text{date}, \text{time})$
 - *Satellite viewing angle* = $F(\text{lat}, \text{lon}, \text{lon}_{\text{sat}})$
 - *Relative azimuth angle* = $F(\text{lat}, \text{lon}, \text{lon}_{\text{sat}}, \text{time}, \text{date})$



Cloud water/ ice mixing ratio → water/ ice contents

Using ideal gas law ($P = \rho RT$),

$$\text{mixing ratio (g/g)} = \frac{\text{water (ice) mass}}{\text{dry air mass}}$$

$$\text{mixing ratio (n}^{\text{th}} \text{ level)} \times \rho_n \text{ (dry air density)} = \text{water (ice) content (g/m}^3\text{)}$$

$$\text{where } \rho_n = \rho_0 \times \frac{P_n}{P_0} \times \frac{T_0}{T_n}, \quad \rho_0 = 1.29 \text{ kg/m}^3,$$

$$P_0 = 1013.25 \text{ hPa}, \quad T_0 = 273.15 \text{ K},$$

$$P_n = P + PB, \quad T_n = (T00 + T) \left(\frac{1000 \text{ hPa}}{P_n} \right)^{-\frac{R}{C_p}}$$

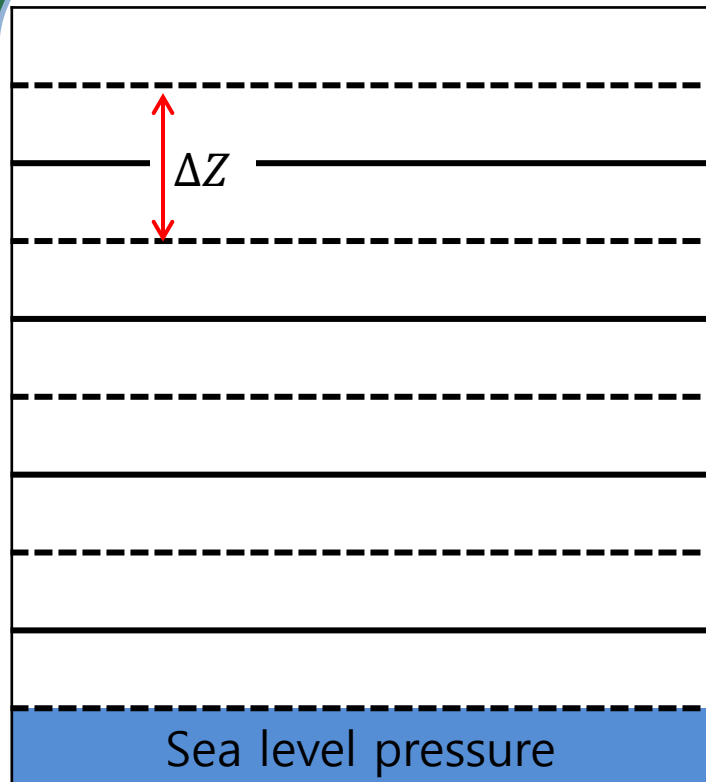
P (perturbation pressure)

T (perturbation potential temperature)

PB (base state pressure)

$T00$ (base state temperature)

Vertical coordinate



← Level of geopotential height

← Level of water or ice contents

liquid or ice water path (g/m^2)

$$= \sum N^{th} \text{ water or ice contents } (g/m^3) \times \Delta Z_N(m)$$

where geopotential height, $Z \equiv \Phi/g_0$

In WRF output,

$$\Phi = PHB(\text{base state potential}) \\ + PH(\text{perturbation geopotential})$$

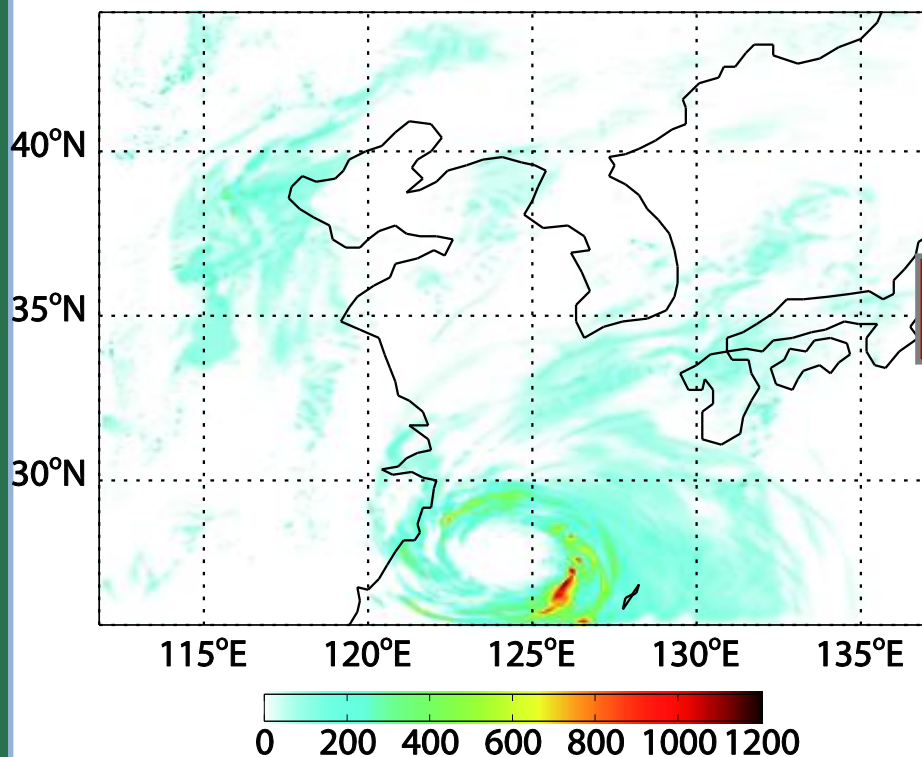
Solid line: unstaggered level

Dashed line: staggered level

An example of simulations

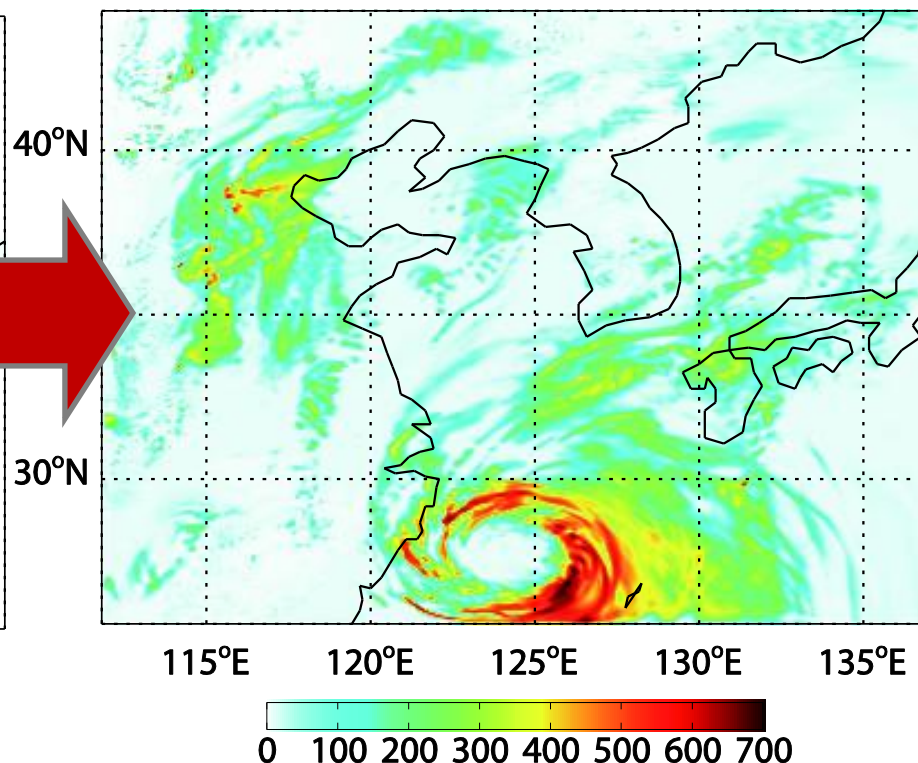
Cloud water path (converted from WRF out)

Cloud total water path (g/m^2)



477-nm Radiance (Total sky)

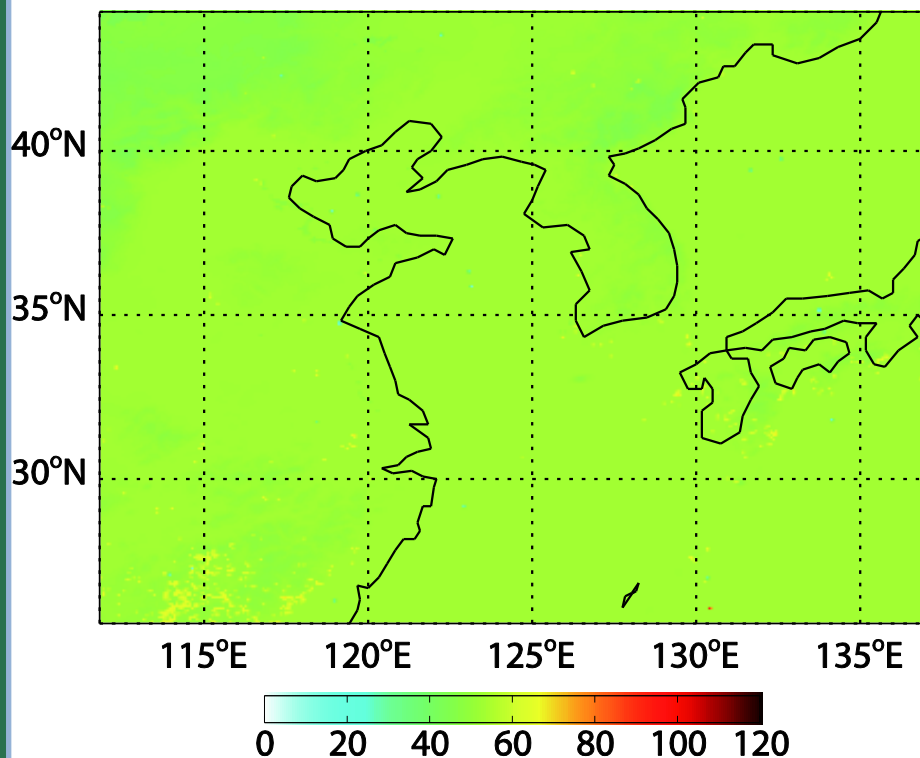
Simulated radiance ($\text{W/m}^2/\text{sr}/\text{micron}$) at 477nm



An example of simulations

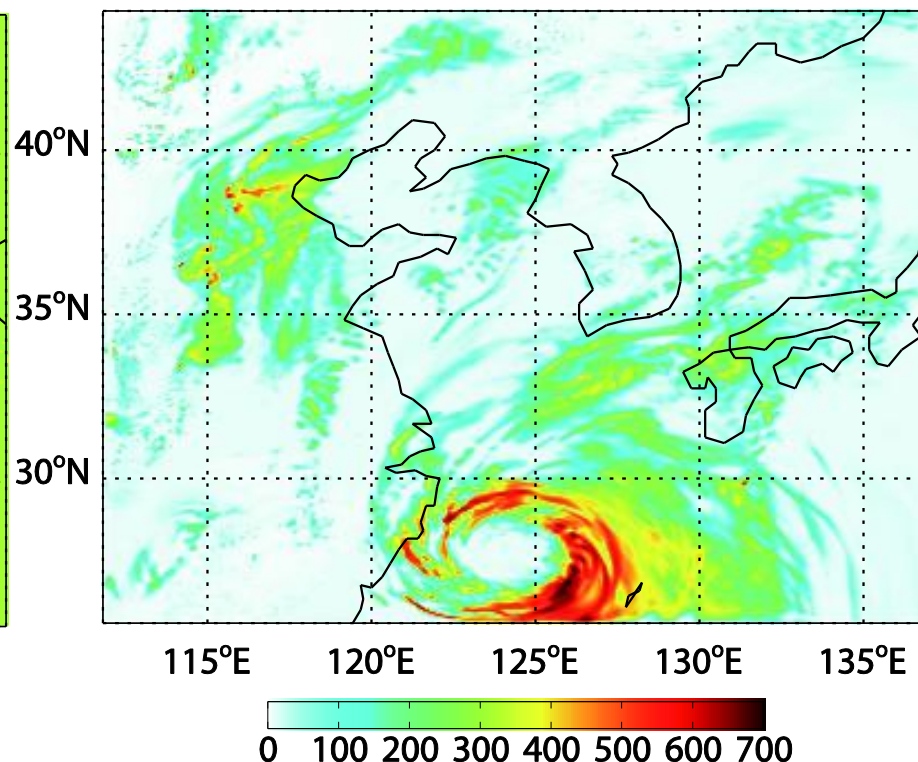
477-nm Radiance (Clear sky)

Simulated radiance ($\text{W/m}^2/\text{sr}/\text{micron}$) at 477nm



477-nm Radiance (Total sky)

Simulated radiance ($\text{W/m}^2/\text{sr}/\text{micron}$) at 477nm



Summary & Future Works

Summary

- ▶ RT simulations from WRF output is ongoing...

Future works include

- ▶ Simulations for the full spectral range of GEMS
- ▶ Simulation of vertically inhomogeneous multi-layer clouds
- ▶ Comparing the retrieved cloud data with the WRF's cloud data
- ▶ Applying the synthetic data to the cloud algorithm

Papers in 2012 supported by this program

- Cho, H., C.-H. Ho, and Y.-S. Choi (2012), The observed variation in cloud-induced longwave radiation in response to sea surface temperature over the Pacific warm pool from MTSAT-1R imagery, *Geophys. Res. Lett.*, 39, L18802.
- Oh, H.-R., C.-H. Ho, and Y.-S. Choi (2012), Comments on "Direct radiative forcing of anthropogenic aerosols over oceans from satellite observation", *Advances in Atmos. Sci.* (*in press*)